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## Multiyear Satellite Total Ozone Column Dimension within West Africa

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### Abstract

Total Ozone Mapping Spectrometer (TOMS\*) mechanisms have been flown on NASA/\*GSFC satellites more than 20 years. They present close to real-time data for Atmospheric Science Research. As elemental of preface energy structured to build up a Lidar base in Nigeria for monitoring the atmospheric ozone and aerosol stratum. The monthly mean TOMS\* total column ozone measurements amid 1978~1999 have been evaluated. The tendency of the total column ozone illustrates a spatial and sequential dissimilarity with cryptogram of the Quasi Biennial Oscillation (QBO) all through this 21-year research epoch; The standards of TOMS\* total ozone vertical, over Nigeria (4-15°N†) is not beyond range of 230~280 Dobson Units, this is reliable with total ozone column data, calculated ever since April 1993 with a total Dobson Spectrophotometer at Lagos (3°22'E†, 6°35'N†), Nigeria.

**Keywords:** Multiyear; Satellite-Lidia-Ozone-Vertical; West-Africa; Measurement; Total.

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## **1. Introduction**

Characterization and parameterization of the silhouette of fluid aerosols, solid particles akin to dust, and atmospheric gases “O<sub>3</sub>, CO<sub>x</sub>” are Imperative for weather prediction, climate replication and environmental observation. Determining the anthropogenic role to the contemporary challenge of ozone depletion could be complicated intimation to the sporadic injection of aerosols from volcanic eruptions, global transport and the interrelation tank of aerosol [1]. Numerous human resources; have detailed the existence of high levels of tropospheric ozone and aerosol in the tropics [2,3].

This phenomenon has profound ecological inferences. Mineral dust may contribute a significant responsibility in affecting the climate by varying the radiation balance in the atmosphere in the course of the radiation scattering and absorption [4,5,6,7]. Mineral dust can furthermore indirectly affect climate through disconcerting cloud nucleation and optical properties [8,9]. Additionally, dust could modify photochemical methods and act as active response surface for reactive gas group in the atmosphere [10,11].

This has been connected to photochemical ozone creation through biomass smoldering [12,13,14,15,16]. It ought to be distinguished that in West Africa (longitudes 22° W~25°E†, latitudes 0°20°N†), the majority of the precipitation and linked thunderstorms is documented throughout the torrential rain months of April to September.

It had been recommended that momentous, circulation, lightning additionally to fires, may be potential donors to ozone creation by producing NO<sub>x</sub> in the middle and upper troposphere [14,17]. Since the atmospheric chemistry over the tropics is greatly influenced by hot rainforests which have immense biomass bustle together with the Savannah where outsized scale bush smoldering incidents are widespread, real-time atmospheric measurements necessitate being undertaken. Atmospheric ozone and aerosol position in West Africa are scant and sporadic. Among the 26 African stop on rout in the World Ozone and Ultraviolet Radiation Data Centre (WOUDC≈) archive, merely one of them, Lagos, Nigeria 3°22'E†, 6°35'N†, 10 m above sea level, is situated in West Africa [18]. The measurements of the entirety ozone vertical ozone with the Dobson Spectrophotometer started at the location in April 1993.

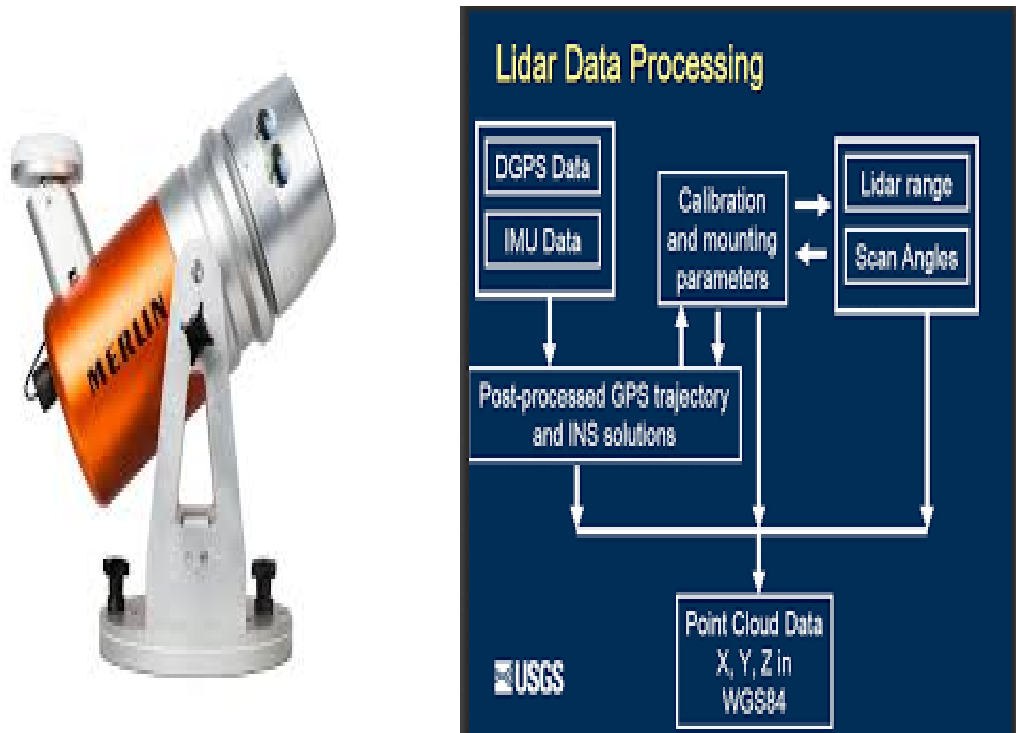
Nigeria, a nation of more than 100 million people with a huge area of land of approximately 100,000 km<sup>2</sup>, is situated amid latitudes 4~15°N† and longitudes 2~18°E†, with sundry climates, arraying from equatorial in the coastal regions, tropical at the center with very dried climates in the northern division neighboring Lake Chad (approximately 16°N†, 15°E†) along with Sahara desert (about 21~25°N†, 11°E†). Concerning the topography, there are southern flatlands amalgamation into innermost tor and upland and peaks in the southeast neighboring Cameroon.

The inadequately studied Cameroon peak which has an altitude of 4101m is situated approximately 11°E† and 5~9°N†, transform the ozone and aerosol levels within West Africa from its recurrent volcanic eruptions. The majority topical volcanic eruption was amid March 28-June 10, 1999 [19]. Ever since Nigeria is a small version of something larger of West African weather and type of weather, observing its ozone and aerosol levels will not

only donate to global atmospheric research nevertheless help in sustainable provincial layout.

Associated learning has been made somewhere else in Africa. Additionally offered seasonal tendency in tropospheric Ozone unhurried in Congo ( $5^{\circ}20'S\ddagger$ ,  $15^{\circ}25'E\ddagger$ ) amid 1983~1986 [14]. They observed that emission of precursor gases from biomass smoldering guide to high ozone standards in the boundary layer particularly through the arid period, in the tropics, ozone is conveyed from the upper troposphere to the subordinate stratosphere owing to the source effect rooted by cumulonimbus thundercloud verticals attendance of excess  $CO_x$  and  $NO_x$  from agricultural actions like biomass smoldering [20,14]. However, Kalicharrun had observed supremacy of Quasi Biennial Oscillation (QBO<sup>§</sup>) even as evaluating TOMS\* Nimbus 7<sup>†</sup> satellite data across South African stop on route [16,17,21].

In this paper, we present the TOMS satellite total column ozone levels within West Africa for the epoch 1978~1999 calculated with the Nimbus 7, Meteor 3<sup>‡</sup> with Earth Probe satellites. This effort disquiets a number of features of QBO<sup>§</sup> in the ozone data within West Africa. Figure1a. Shows the Lidar instrument while figure1b illustrates Lidar data processing; this learning can be measured a standard of value for our lasting goal of scheming a Lidar location for observing tropospheric and stratospheric ozone and aerosol stage in the country Nigeria.



**Figure 1a:** demonstrated Lidar while figure 1b illustrates Lidar data processing. (Courtesy Atmos)

The positioned, “ $5^{\circ}30'N\ddagger$ ,  $7^{\circ}12'E\ddagger$ , and 35 m above sea level”, at Imo State University, Owerri could comprise of a “degree of difference assimilation Lidar” “DDAL” scheme and Rayfeigh-Mie Lidar [22].

### **1.1 Definition of problem**

To comprehend and elucidate total ozone column dimension within West Africa by means of satellite. To elucidate the importance of TOMS\* data pro Atmospheric Science Research for the advance atmospheric scientific knowledge. To evaluate the initial assessment quandary of total ozone by means of a chemical transport model which is probable to be the basis of facilitating middle of series beyond several -year's prediction of the surface UV-B radiation. Nevertheless, to understand that the absence of vertical information, the absorption of ozone columns has only petite impact on the mold of the vertical ozone silhouette, which is principally gritty by the transport. And to evaluate the convolution of natural and anthropogenic chemistry and physical introduction of elementals and there effects Ozone layer.

### **1.2 Objective of Study**

The aim of this study to provide long-term series of high-level absorbed TOM's\* data to scientific clients. The outcome of this research will give us exhaustive quality estimates and validation data sets of the TOM'S products, based on an unmitigated set of ground based and ozone observations, and the error data acquired from the data absorption. To understand data sets that will be deal with productivity from global chemistry-transport molds to advance their modeling capabilities of contemporary and potential revolutionize of ozone and chemical dynamic greenhouse gases that comprises:

- Technique of coalesced dynamical knowledge with the chemical observations of TOMS\*, consequential in various dimensional of ozone and NO<sub>2</sub> statistics sets.
- To mutual comprehend parameters affecting the precision of ozone sonde silhouette that could be practical to all accessible ozone sonde measurements for the TOMS\* epoch.
- To adopt method that will endow with approximation of the stratospheric and tropospheric NO<sub>2</sub> columns.

### **1.3 Significance of the Study**

The importance of this study is on data absorption characteristics and the analysis domino effect.

- The fundamental motivation is to evaluate the presence of pandemonium within the appropriate time sequence.
- Generate unrestricted knowledge in sequence to the hazard foundation to climate change
- To understand the causes of Ozone depletion which is as results of complimentary radical catalysts, comprises:
  - ✚ nitric oxide (NO),

✚ nitrous oxide (N<sub>2</sub>O),

✚ hydroxyl (OH),

✚ atomic chlorine (Cl), and

✚ atomic bromine (Br).

- Though there are natural foundations for all of these species, the clusters of chlorine and bromine augmented distinctly in modern decades for the reason that of the liberation of huge quantities of anthropogenic organohalogen compounds, principally chlorofluorocarbons (CFCs) and bromo\*fluorocarbon [27].
- These highly established compounds are able of in existence the rise to the stratosphere, where Cl and Br radicals are liberated by the accomplishment of ultraviolet light.
- To understand the hazed of breakdown of ozone in the stratosphere its consequences in abridged absorption of ultraviolet radiation.
- Therefore, unabsorbed and hazardous ultraviolet radiation is able to arrive at the Earth's surface at a higher intensity. Ozone heights have plummeted by a worldwide average of approximately 4 percent ever since the late 1970s. For roughly 5 percent of the Earth's surface, in the regions of the north and south poles, to a large extent seasonal decline have been seen, and are depicted as «ozone<sup>def</sup> holes» [26]. The breakthrough of the yearly depletion of ozone beyond the Antarctic was first publicized by Joe Farman, Brain Gardiner and Jonathan Shanklin, in a paper which emerged in [28].

#### ***1.4 Limitation of Study***

The precincts of this study comprise the subsequent:

- ❖ The convolution in gases specimen into the atmosphere, compelled the researcher to specimen some hard metals in the atmosphere.
- ❖ Conversely, for the reason that of the weak outline silhouette effects beneath for the most part observing circumstances for wavelengths longer than 315 nm and the non insignificant ambiguity in measurements by space<sup>def</sup>borne instruments, there is more often than not limited information about the vertical circulation limited in these measurements, debarring exhaustive silhouette retrievals.
- ❖ Contemporary data of the archetypal vertical SO<sub>2</sub> circulations for mutually natural sources and anthropogenic is very limited. Which implies that SO<sub>2</sub> from industrial air pollution over and above oxidation of natural matter is likely to be cramped to the planetary boundary layer (PBL<sup>m</sup>), despite the fact that SO<sub>2</sub> from unrestrained eruptions or degassing of volcanoes is likely to broaden within a constricted layer at a height comparable to the altitude of the foundations, and SO<sub>2</sub> through explosive

volcanic eruptions could be infused into the upper troposphere or lower stratosphere.

## **2. Hypothetical Milieu**

A German scientist by name Christian Friedrich Schönbein was the first to find ozone in 1839, identification it subsequent to the Greek word ὄζειν denotation «to smell», as recommended by his associate W. Vischer, professor of greek in Basel [38]. However, some French physicists Charles Fabry and Henri Buisson discovered the ozone layer in the stratosphere with its magnitude to absorb the sun's ultra-violet emissions in 1913 [39]. Furthermore, Ozone depicts approximately merely 0.00001% of the atmosphere by volume according to which means that there are approximately two ozone molecules for each 10 million air molecules [40]. Conversely, atmospheric ozone plays a fundamental role that contradicts this small relative size. Ozone absorbs for the most part of the UV-B radiation as well as a foremost part of the UV-A radiation coming from the sun, thereby permitting only a petite fraction to reach the Earth's surface.

Ozone is discovered in two expanse of the Earth's atmosphere: roughly 90% of the ozone exists in the stratosphere, approximately 10-50 km beyond the Earth's surface; the residual 10% of ozone is at the troposphere, which broadens itself to an altitude of 10 km. However the foundation of both is dissimilar. At the stratosphere, ozone is produced by oxygen photolysis, tempt by UV-C solar radiation. Upon entering the atmosphere, solar radiation could basically be dispersed, generating diffused radiation, or it can interrelate selectively, primarily to modify in the chemical composition of some elements of the atmosphere.

Ozone plays a double function at the stratosphere; is advantageous for an assortment of forms of life on Earth. While on the other hand it defends the Earth's surface against a great deal of the UV-B and UV-A radiation, also on the other hand it acts a structural function in atmospheric temperature. Consequently the formation and devastation sequence active response of natural ozone in the ozone layer are recognized as Chapman reactions [41]. Stratospheric ozone is damaged by photo-dissociation, by engrossing UV-A and UV-B radiations. Subsequently the photon's energy is used to smash the link in ozone molecules and the energy overload is transformed into kinetic energy of the oxygen molecules along with the atomic oxygen generated and released in the dissociation. The kinetic energy of the photo-dissociation consequences in an increase "↑se" of the gaseous exciting, and consequently generates a momentous temperature gradient at the stratosphere. Come at the troposphere, ozone formation approach from effect generated by combustion, but a petite fraction is generated naturally by the atmospheric electrical discharges. However, the ozone in the surrounding area of the surface of the Earth in get in touch with life forms has the chattels to act in response strappingly with a digit of molecules and becomes chemically lethal to living organisms.

About hundreds of millions of years, the stratospheric ozone concentration continue stable. However, according to forecasted that ozone will be damaged by the act of increasingly accumulated chlorofluorocarbons (CFC) discharged into the atmosphere [42]. Chlorine in these compounds operates as the catalyst in the obliteration of the ozone layer. The obliteration of the ozone layer did not occur only over the Antarctic; nevertheless in 2003, the scientific evaluation of ozone depletion exposed a drop in ozone levels of roughly 4% per decade in mid latitudes [27]. However, among others, report stabilization in stratospheric ozone ever since 1997 [43].

## 2.1 Methodology and Statistics

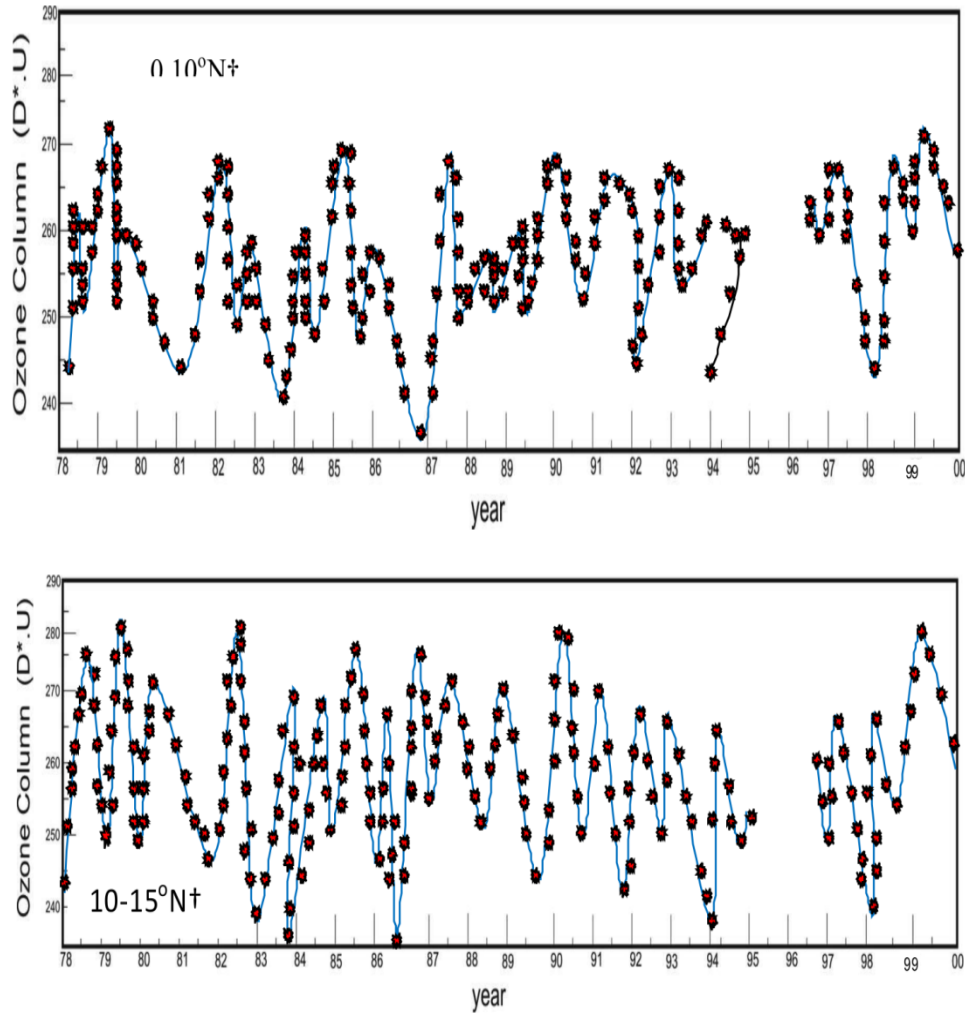
Total Ozone Mapping Spectrometer (TOMS\*) mechanisms flown on four NASA/\*GSFC spacecraft: Nimbus 7 by November 1978~May 1993; Meteor 3 in August 1991~December 1994, \*ADEOS in September 1996~June 1997 and Earth Probe in July 1996 current. In this research, the Nimbus 7, Meteor 3 and Earth Probe TOMS\* regional monthly mean entirety vertical ozone data for the years 1978-1999 were evaluated by means of Grid Analysis and Display System (GRADS\*) software (<http://www.grads.iges.org>) [23]. The \*ADEOS statistics that totally extend beyond the Earth Probe was not adapted. At that point no TOMS statistics pro the year 1995 and amid January/July 1996.

The TOMS\* data for daily are gridded indicated  $\Delta^1$ Lat. degree regions by  $\Delta^1$ .26 Long., degree sectors. Latitude proceed from -90 degrees South Pole† to 0 degrees equator to +90 degrees North Pole† in  $\Delta^1$  degree short way consequently at that point are 182 west latitude region. Alternatively, the longitudes move starting -180 west longitudes to 0 Greenwich; ~England to +182 east longitudes in  $\Delta^1$ .26 degree short, which implies approximately 289 longitude regions. Other niceties with reference to the TOMS\* instruments like orbital typical, mechanism and quantifying methods can be established on the NASA TOMS\* web page [24].

For purpose of this exertion; West Africa is taken as the area of Africa amid latitudes 0-25°North† and longitude 25°West† to 25°East†. This was concluded to capture the modulators of the neighborhood weather and climate across the region into detailed examination. They comprise Cameroon mountain approximately 12°E†5-8°N† which is a acknowledged foundation of volcanic aerosols, tropical forest linked biomass smoldering in the Sahel province, gases occurring through fossil fuel smoldering with Ultraviolet (UV) engrossing aerosols through arid regions close to Lake Chad «approximately 16°E†. 15°N†) while Sahara desert close to 21°N†. Such arid regions hold loads of fine particulate matter that be able to effortlessly conveyed by means of winds, presently as well about large digits of oil wells in Nigeria with allied gas flares ever since the belatedly 1960's. Concerning the climate of the knowledge region, the majority of the precipitation is documented amid the months of April/October, whereas the major arid months are October and March, The climax gear of biomass smoldering are amid Dec.~ March whereas the lightning discharge is obvious all through the damp torrential rain months. In advance workers like they had observed that biomass smoldering and lightening discharge influence the tropospheric ozone stages in the tropics [13,15,3,4,21].

## 3. Regional Monthly Mean Ozone Point within West Africa amid 1978-1999

The precinct mean TOMS\* entirety vertical ozone statistics for the years 1978 to 1999 across West Africa 0~10°N†. 10~15°N†, 10~20°N†, 15~25°N† latitude groups are accessible in figure 3. 1 and figure 3. 2. No dimensions were accessible for the months of June 1978, March, June, July and October 1993 and January, February, August, September and December 1994. Although in 1995, as well as from January to July 1996.

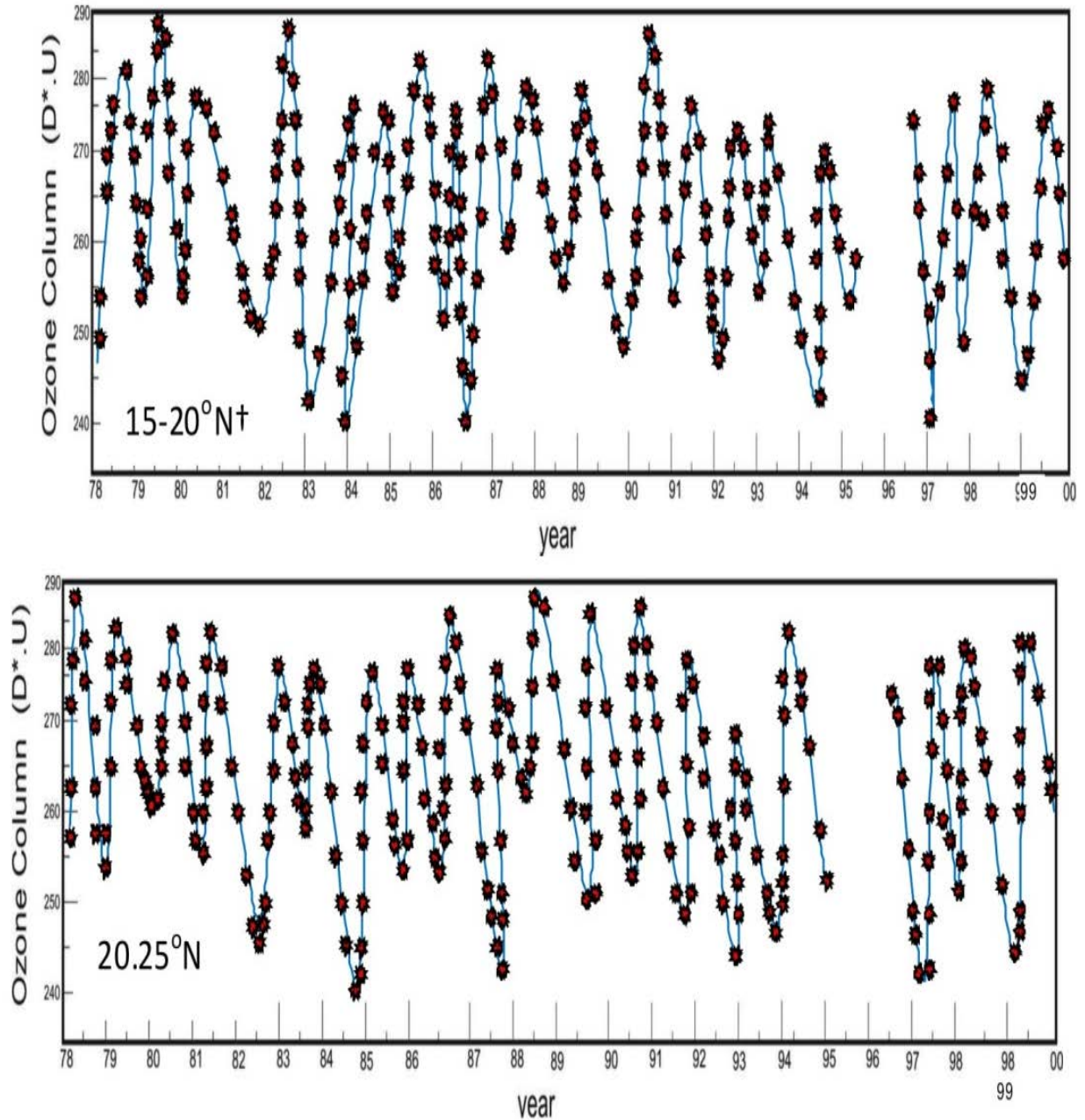


**Figure 3.1:** prescient mean TOMS\* Total vertical ozone level over West Africa amid 1978-1999 in the months of Jan.-Dec. for latitude 0-10°N† and 10-15°N†.

The least standards of the TOMS\* ozone data were illustrated in figure 1 and figure 2 and were documented in the months of December to February when the lesser atmosphere over the area included dust particles from the Sahara desert. This epoch is also typified by biomass smoldering, which extensively influence the photochemical creation of ozone in the lesser ambience [15,21]. This is marked difference with the higher standards all through the wet period April-September, with the climax sum experiential amid July and September, while storm with thunder and lightning as well as dynamical circulation, are improved.

An uncomplicated assessment of TOMS\* ozone statistics utmost and lowest demonstrates that similar standards happen in approximately 24-months. This is supplementary manifest in 0~10°N† 10~15°N† groups as illustrated in figure 1. We additionally carried out the Quasi biennial Oscillation “QBO<sup>δ</sup>” autograph by plotting the distinction amid the total ozone vertical and the bi-monthly climatologically mean illustrated in (figure 3 and figure 4). In the 0-10°N† group of West Africa, the “QBO<sup>δ</sup>” distinctive part are quite obvious beside the entire statistics set. The QBO<sup>δ</sup> appears tranquil to be present in 10-15°N† group, except with an abridged amplitude.

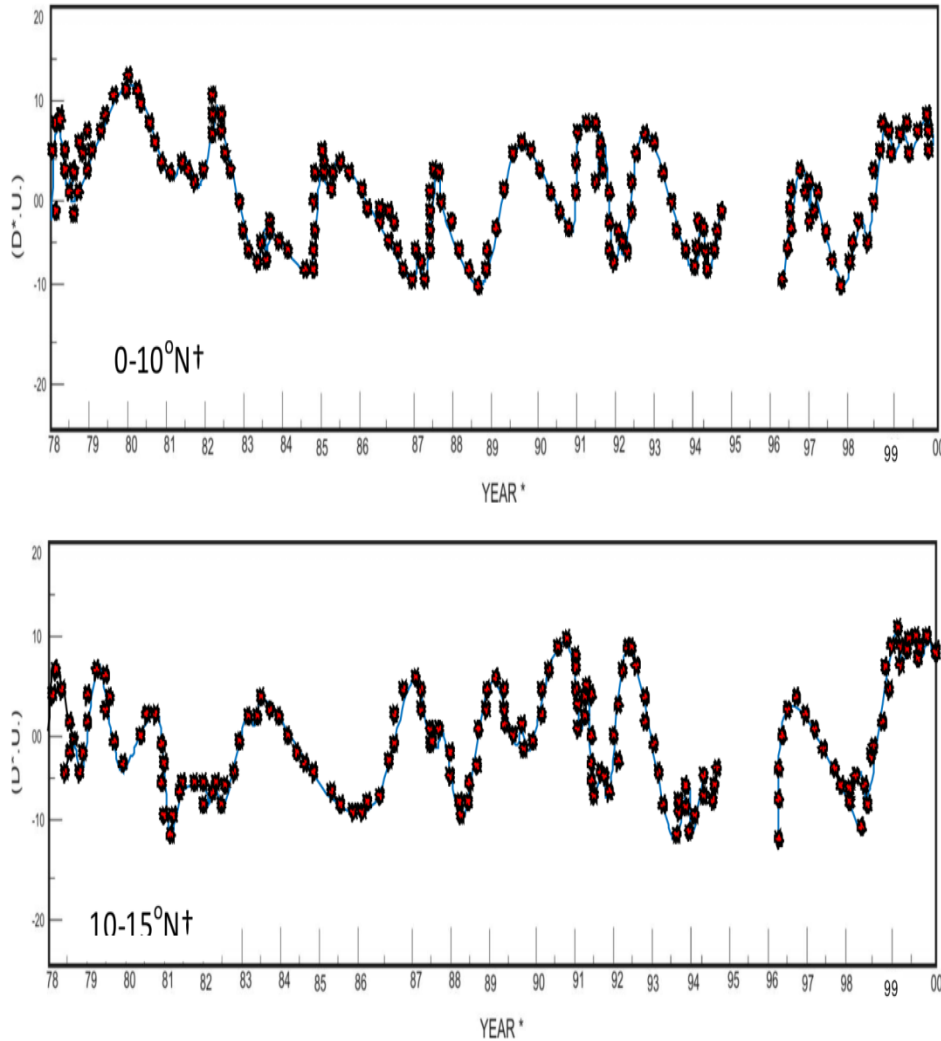




**Figure 3.2: figure 3.1:** illustrates for latitudes 15-20°N† and 20-25°N† correspondingly.

**Table 1** illustrates regional data of TOMS\* satellite total ozone vertical depth measurements over West Africa with reference to the 21-years period; we detailed the annual average standards, utmost and lowest, with standard deviations<sup>δ</sup>. Most of the utmost are in July~September epoch, when biomass smoldering bustle is lesser than lightning results. Amid 0-10°N†, the below average worth of the total ozone was 246.6 DU February of 1987 \*El Nino year. The 0-10°N† 1987 statistics as well have the highest worth of standard deviation<sup>δ</sup>. The statistics for the other \*El Nino years (1982 as well as 1983) illustrates comparable features. Usually the data demonstrate that the total ozone Vertical augmented from the equator inwards. The standard deviation<sup>δ</sup> of TOMS\* satellite ozone size are higher amid latitudes 10-20°N† where the results of biomass smoldering can be

further well-defined.

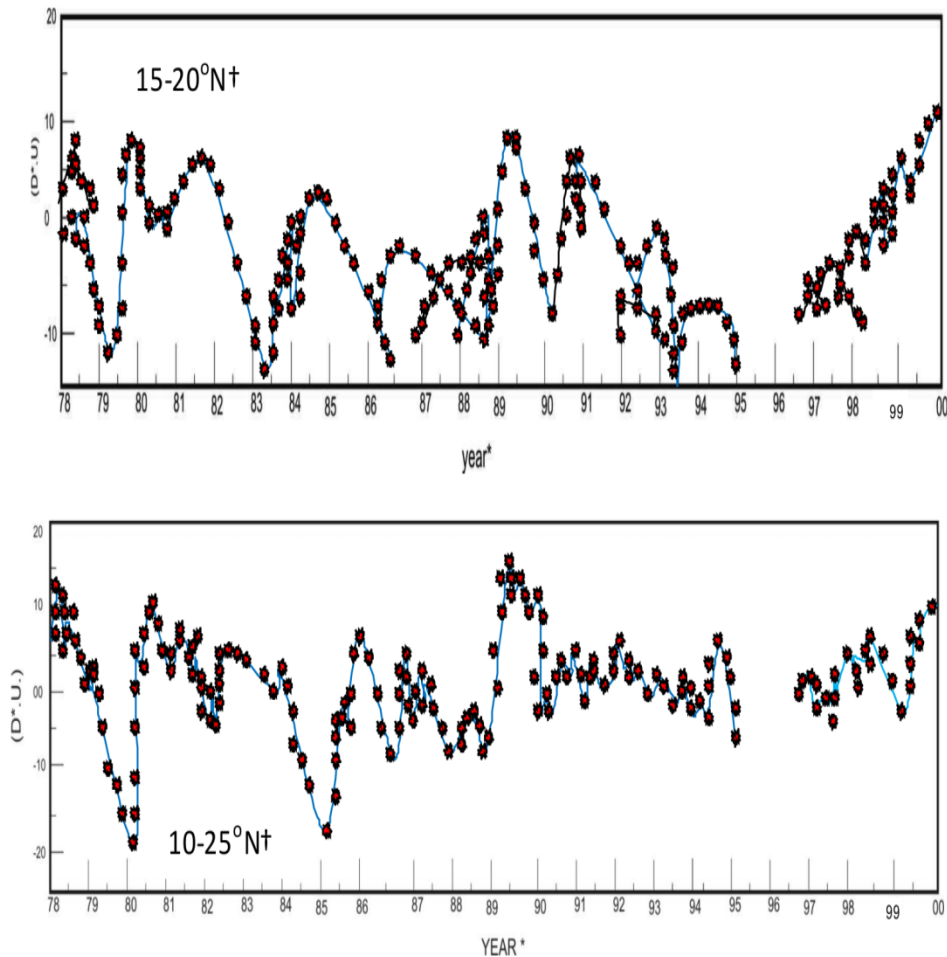


**Figure 3.4:** Illustrates divergence amid the Zonal mean and monthly average (1978-1999) TOMS total column ozone levels over West Africa between 1978 and 1999 in the months of January to December for latitude 0-10°N† and 10-15°N†.

#### 4. Discussion and Conclusion annotations

In this introduction exertion, we evaluated the zonal mean TOMS total column ozone statistics précised by three of the NASA/GSFC satellites (Nimbus-7, Meteor-3 and Earth Probe amid January 1978 also December 1999). The hiatus is long sufficient to make some notes concerning the Quasi Biennial Oscillation tendency over West Africa, The dry period (January-March), which also inscription the north nearly all gradient of the Inter tropical Convergence Zone (ITCZ), documented the lowest stage of ozone, whereas throughout the damp period, the highest sum of total ozone was experiential, The West African ozone levels could be changed by lightning discharges, further obvious augment emerge in August, through the south nearly all decline of the ITCZ [15,17,18]. A comprehensive assessment of photochemical progression complicated increasing reaction rate

ozone obliteration in the lesser atmosphere will almost certainly shed further light on this. Some distinctive part of \*El Nino Southern Oscillation (1982, 1983 and 1987) obviously influence TOMS satellite dimensions, the responsibility of \*El Nino Southern oscillation observable fact, principally on the tropospheric ozone quantity, needs necessitate to be scrutinized by means of height determined measurements (explicitly Lidar). The TOMS notes in the 10-15°N† band consent with the dimensions of a Dobson spectrophotometer located at Lagos, Nigeria.



**Figure 3.5: figure 3:** illustrates latitude 10-15° N† and 15-20°N†

Ever since long-term objective of this exertion is the establishment of a lidar location for observing ozone and aerosols in the troposphere and stratosphere over Nigeria, the domino effect accessible in this paper and individuals of aforementioned workers over Africa will be a advance in that consideration [15,21]. At the same time as greatly welcome the exertion of the NASA/Goddard Space Flight Center (Code 916\*) for incessant hard work at sustaining the TOMS agenda ever since its inception over 21 years ago; the need for proliferate the current complex of ozone Sonde as well as lidar stations world-wide, principally in the tropics, needs to be painstaking for a better consideration of the troposphere and stratosphere.

**Table 1:** shows regional data of the TOMS\*, Satellite Total Vertical Ozone Dimensions across West Africa amid 1978 and 1999. The lowest, utmost and mean standards of the Total Vertical Ozone statistics are in Dobson Units, is standard deviation<sup>δ</sup> S<sup>δ</sup>d

Year	0.10° N†				10-15°N†				10-20°N†				15-25°N†			
	Min.	Max.	Mean	S <sup>δ</sup> d	Min.	Max.	Mean	S <sup>δ</sup> d	Min.	Max.	Mean	S <sup>δ</sup> d	Min.	Max.	Mean	S <sup>δ</sup> d
1978	247.1	271.1	260.2	7.8	245.3	279.5	262.8	12.1	247.1	284.2	266.6	14.2	247.3	282.5	270.7	14.2
1979	247.2	271.2	260.1	7.8	245.4	279.5	262.8	12.0	246.1	284.2	266.7	14.1	248.0	288.0	270.8	14.3
1980	256.3	284.9	269.9	10.2	247.8	286.6	267.8	13.8	241.9	286.2	265.5	15.9	243.5	285.8	266.7	15.3
1981	247.0	269.6	260.7	6.6	247.6	276.3	263.5	9.9	248.6	281.3	266.0	11.6	253.2	282.7	270.9	11.3
1982	243.5	281.4	263.6	12.8	242.3	283.1	264.4	14.1	241.9	281.6	264.7	14.1	242.0	283.7	268.9	14.3
1983	244.6	267.3	257.3	7.1	240.6	272.8	257.7	11.4	237.9	275.7	259.5	13.6	239.2	277.4	263.1	13.5
1984	240.0	266.8	254.1	8.4	239.5	270.9	256.8	10.5	240.0	275.7	259.7	12.8	235.8	279.5	264.3	14.5
1985	249.6	279.6	267.4	9.6	247.2	281.8	265.8	12.8	238.4	279.9	262.8	14.8	232.2	278.7	261.8	16.0
1986	242.7	262.8	254.3	6.9	241.2	270.0	257.1	9.8	243.1	273.5	261.3	10.6	247.0	278.5	266.1	10.8
1987	234.9	279.5	258.6	14.5	235.6	277.8	259.7	13.8	239.1	275.9	260.3	13.6	238.0	282.6	263.6	15.4
1988	249.4	266.2	259.1	5.5	248.5	272.0	260.3	8.8	241.2	275.7	261.3	11.8	237.8	277.2	262.4	12.9
1989	250.3	270.0	260.4	6.5	250.0	276.3	264.6	10.5	240.7	283.6	269.4	13.1	246.7	289.5	274.1	14.3
1990	255.3	283.9	271.3	10.1	247.0	285.1	269.4	13.0	240.4	284.0	266.8	14.4	241.7	289.4	267.6	14.1
1991	244.1	269.9	258.1	8.1	241.3	279.2	262.0	13.0	237.7	283.0	264.8	15.3	244.8	284.4	268.4	15.4
1992	242.3	277.6	261.6	10.7	238.1	276.3	260.5	11.8	240.5	276.1	259.6	12.6	237.6	285.4	262.6	14.6
1993	238.4	263.8	256.2	8.7	238.8	266.7	255.0	9.2	248.3	269.2	254.4	11.4	239.0	270.2	255.5	13.0
1994	250.3	256.8	256.2	5.7	251.1	266.7	258.2	5.7	241.2	273.9	261.3	8.4	245.5	279.6	265.9	11.6
1996	255.0	263.8	260.3	3.4	247.4	268.5	260.0	8.1	240.1	274.3	260.7	12.3	240.0	275.8	260.9	13.3
1997	243.9	274.2	265.4	9.3	243.5	276.6	263.5	11.4	239.3	278.0	261.8	13.3	237.8	279.5	262.5	13.7
1998	239.7	275.5	260.6	11.1	237.5	276.8	261.5	12.3	245.1	281.5	264.1	13.9	245.3	284.8	267.9	14.9
1999	256.4	282.0	271.6	8.5	253.7	283.3	270.0	10.8	241.9	283.7	268.2	13.5	241.9	284.7	268.5	15.0
<b>Mean</b>	<b>246.6</b>	<b>272.2</b>	<b>261.2</b>	<b>10.9</b>	<b>243.7</b>	<b>276.5</b>	<b>262.9</b>	<b>15.9</b>	<b>242.0</b>	<b>266.0</b>	<b>263.3</b>	<b>13.9</b>	<b>242.2</b>	<b>280.1</b>	<b>265.9</b>	<b>13.9</b>

The wave†reminiscent of nature of every day ozone concentration illustrates a periodical sequence comprising wavelength of approximately 741 Julian days «preliminary as of 1<sup>st</sup> January, 2001», in the sagacity that merely half of the cycle is accomplished in a year while the residual part is accomplished in the following year.

The sequences begin with a smallest throughout the period of dry season and increasingly presuppose an utmost throughout the period of wet period. Lowest amount and utmost ozone concentrations within apiece station are found to transpire principally in January and June correspondingly. Nevertheless in West Africa the epoch of dry season begins in October and lengthens to March of the subsequent year while period of wet season begins in April and ends in September.

This corresponds to a lag of two months amid the onset of a season and least or utmost ozone concentration as the case could be. The coordination of ozone autocorrelation turn out to be decoupled subsequent to a lag of 2 months indicated by non±persistence in the autocorrelation as the lag increases “†s”, ordering to the squat memory of ozone scheming factors.

Vigorous weather makes certain sturdy mixing, forceful transport and raid/substitute of ozone owing to interaction of stratospheric and tropospheric elementals. Comprehensive discussion of West African wind method entails mentioning the dynamics and functions of tropical Easterly† Jet (TEJ $\mathbb{W}$ ), African Easterly† Jet (AEJ $\mathbb{W}$ ), Easterly† Waves (EW $\mathbb{W}$ ) and Inter†tropical discontinuity (ITD $\mathbb{W}$ ), for further on this topic concerned reader is directed to, for instance [29,30]. Conversely, the captivating facet of these wind patterns is that they interrelate to generate and/or strengthen two foremost seasonal winds in West Africa, that is to say: the moist south western and the dry north eastern winds, such relations has been explored to elucidate excellent and poor rainfall freedom in West Africa according to [29,31].

The involvement is that weather actions unaccompanied can clarify about 63% of ozone circulation in West Africa. Just less than 40% are owing to the effect of biomass burning and others. The weather action in this paper refers to the general effects of interaction amid stratospheric and tropospheric wave action which in the case West Africa obviously evident into two different seasons as stated above.

As precipitation is most effective meteorological stricture that delineates the periods, we have used precipitation here. Therefore, ozone loss or gain can be credited to dynamical interaction amid stratospheric and tropospheric atmosphere determined by weather bustle [32,34].

In the months of July-September as the utmost frequency of precipitation is documented in the area, the prevalence of lightning' and linked unconstrained NO<sub>x</sub> gases is utmost across West Africa. The standards of the TOMS satellite total column ozone measurements were lowest all through the months of October-February, when the prevalence of biomass smoldering and circulatory atmospheric recent occurring from lightning discharges is below average. According to the circulative regions, which are generally allied with biomass smoldering, persuade vertical amalgamation thus disturbing the ozone cluster in the lower atmosphere [25].

Furthermore, the measurements carried out at Brazzaville, Congo had endowed with substantiation for huge-scale photochemical ozone creation in West Africa [4]. These domino effects concur with individuals of

[15,26,4].

Though considerable negative correlation of approximately - 1.00 was experiential amid the precipitation design circulation and total ozone percentile inconsistency over the five zones in West Africa on spatial foundation, however, in provisos of incidence, the epoch of utmost ozone cluster\*ation agree with the crest tropical summer precipitation. Tropical summer precipitation over the West African area crests amid June and September.

This assessment might probably be credited to diminution in the potency of the further tropical suction pump (FTSP $\pm$ ) movement accountable for the transportation of ozone through the tropical stratosphere into the middle and high latitudinal area. The FTSP $\pm$  is a phenomenon, by means of which the more than tropical stratosphere and mesosphere in the course of having social significance eddy effects operate globally on the tropical stratosphere as a fluid $\ddagger$ dynamical suction pump [44,45,46]. Consequently it may be indicate that there is mutually connecting amid decline in the strength of the FTSP $\pm$  and ozone distribution all through the tropical precipitation season. Diverse magnitude of sunlight established at the surface of the earth (known as solar insolation) coerces the weather and the experiential winds on the earth. Similarly, temperature modification crosswise earth's surfaces are directly connected to the speed and course of the winds, both at the surface and at diverse heights [47].

## **5. Conclusion**

Apparently, this study radiates light on the ozone spatial and sequential distribution over West Africa area. The function of precipitation generating mechanism has been stressed. Ozone distribution is typically proscribed by precipitation generating device such that ozone accumulation go after synchronously the relocation of wind patterns. The moist south westerly deriving from Atlantic Ocean flows to improve ozone accretion as ozone utmost concentration arise when it takes over the area. On the other hand, north eastern complement reduces ozone concentration in that way producing lowest ozone for the duration of the dry season. In regard to the two months lag amid southerly and northerly precipitation, this study has shown that same lag happened amid south (lower $\pm$  latitude) and north (higher $\pm$  latitude) monthly ozone utmost.

Oscillating biennial characteristic demonstrated by ozone interannual tendency is well thought-out as the route of the upper level atmospheric distribution explained to as quasi biennial oscillation (QBO). This distribution performs to “ $\uparrow$ se” the stratosphere  $\S$  troposphere ozone substitute. Bourke and Lome correspondingly evaluated to period of weakened QBO. The QBO influence the global stratospheric distribution, and extends to  $\pm 22^\circ$  north $\dagger$  and south $\dagger$  with an amplitude of approximately 12 m/s [35]. It influences an assortment of extra $\pm$ tropical phenomena as well as the vigor and permanence of the winter time polar eddy, and the circulation of ozone and other gases [36]. The QBO is motivated by the dissipation of an assortment of equatorial waves, that are primarily mandatory by profound cumulus convection (weather actions $\equiv$ ) in the tropics [37,34].

## 6. Recommendation

Ground-based networks: Brewer and Dobson spectrophotometers, DOAS UV-visible and FTIR spectrometers, lidars, ozonesondes, millimeterwave radiometers, and ground-level *in situ* size) comprise the primary foundation of correlative annotations for the fundamental validation of satellite data. Such networks ought to be upholding, even unlimited, to cover a wider array of atmospheric nations, states and areas of interest for the advancement of Atmospheric sciences. The recent “↓se” in the digit of stations document given information data to WOUDC could become a worry if this decline of facilities prolong. The deployment of essential facilities in the tropics and the Southern Hemisphere is very optimistic. For some class and instruments, advance exertion is essential to advance the station to station homogeneity of networks, and to strengthen long\*term data accounts, particularly in view of data evaluations addressing the links amid atmospheric composition modifications and climate change.

There are numerous matters that linger on the predictable ozone recuperation from the influence of ozone-depleting material (ODMs≠). Particularly, how do ozone depletion and climate change interrelate? Modern research disclosed that ozone depletion has influenced tropospheric climate. Furthermore, it is becoming lucid that greenhouse gases (GHGs) are varying the stratosphere; the cooling of the upper stratosphere by GHGs is anticipated to exceed 6K amid the years 2000 and 2100, obliging lasting annotations of mutually ozone and temperature in the stratosphere.

The capacity to forecast upcoming ozone actions necessitate further advancement in the quantification of the functions of chemical and dynamical progression accountable for ozone assembly, transport, loss, and circulation, and their individual uncertainties. The advancement of practical scenarios of the upcoming abundances of anthropogenic and biogenic trace gases in the stratosphere and troposphere is essential, principally with regard to a changing climate.

Reproduction as of the 2010 scientific evaluation of Ozone depletion signify potential ↑ses of UV levels in the tropics, but ↓ses at mid≡and≡high latitudes owing to ozone changes. The 2010 details of the environmental effects assessment panel (EEAP) accomplished that research on the impacts of ↑ses in UV radiation consequential from stratospheric ozone depletion has significantly advanced the understanding of the progression by which revolutionize in UV radiation influence a array of organisms and procedures. Pro humans, this creates the risk of further skin cancer in the tropics, but also slightly ↑ses the risk of UV doses that are too low for the assembly of adequate Vitamin D at mid≡high latitudes.

Latest research has tinted the interactions amid the diverse effects of altering UV radiation owing to ozone depletion and the effects of climate change. These interactions could lead to feedbacks into climate change e.g., alteration of carbon cycling in terrestrial as well as aquatic ecosystems, nevertheless this yet inadequately defined.

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### **Conflict of Interest**

The authors assert no clash of interest.

### **References**

- [1]. Barnes.J.E. and D.L. Hofmann (1997): Lidar measurements of Stratospheric aerosol over Mauna Loa observatory. *Geophys. Res, Lett*, 24 (15), 1923-1926.
- [2]. Fishman, J., Watson, C.E., Larsen, J.C. and Logan, J.A. (1990) Distribution of tropospheric ozone determined from satellite data. *Journal of Geophysical Research* 95: doi: 10.1029/89JD02784. issn: 0148-0227
- [3]. Andreae M.O. and W.A. Jaeschke, 1992: Exchange of Sulphur between biosphere and atmosphere over temperate and tropical regions, in: *Sulphur cycling on the continents: Wetlands, Terrestrial Ecosystems, and Associated Water Bodies*. SCOPE 48, R.W. Howarth, J.W.B. Stewart and M.V. Ivanov (eds), Wiley, Chichester, 27-61.
- [4]. Tegen I., Hollrig P., Chin M., Fung I., Jacob D. & Penner J., 1997, Contribution of different aerosol species to the global aerosol extinction optical thickness: Estimates from model results, *J. Geophys. Res.*102:23895– 23916.
- [5]. Haywood. J; Boucher. O; First published: November 2000 Full publication history; DOI: 10.1029/1999RG000078View/save citation
- [6]. Harrison and Aplin 2001; Eichkorn *et al.* ... doi:10.1088/1748-9326/8/1/015026 S. Eichkorn, S. Wilhelm, H. Aufmhoff, K.H. Wohlfrom, F. Arnold, References
- [7]. Sokolik, I.N., Winker, D.M., Bergametti, G., Gillette, D.A., Carmichael, G., Kaufman, Y.J., Gomes, L., Schuetz, L. and Penner, J.E. (2001). Introduction to special section: Outstanding problems in quantifying the radiative impacts of mineral dust. *Journal of Geophysical Research* 106: doi: 10.1029/2000JD900498. issn: 0148-0227.



- [8]. Levin Z., Ganor E. & Gladstein V., 1996, the effects of desert particles coated with sulfate on rain formation in the eastern Mediterranean, *J. Appl. Meteorol.* 35:1511–1523.
- [9]. Wurzler, S., T. G. Reisin, and Z. Levin, Modification of mineral dust particles by cloud processing and subsequent effects on drop size distributions, *J. Geophys. Res.*, 105, 4501–4512, 2000
- [10]. Dickerson, R. R., Kondragunta S., Stenchikov G., Civerolo K. L., Doddridge B. G. & Holben B. N., 1997, The impact of aerosols on solar ultraviolet radiation and photochemical smog, *Science* 278: 827–830.
- [11]. Dentener F. J., Carmichael G. R., Zhang Y., Lelieveld J. & Crutzen P. J., 1996, Role of mineral aerosol as a reactive
- [12]. Fishman, J. and Larsen, J.C. (1987). Distribution of total ozone and stratospheric ozone in the tropics: Implications for the distribution of tropospheric ozone. *Journal of Geophysical Research* 92: doi: 10.1029/JD092iD06p06627. issn: 0148-0227 Andreae, M.O., Talbot, R.W., Andreae, T.W. and M.C.,
- [13]. Andreae, M.O et al, Biomass burning emission and association haze layer over Amazonia, *J. Geophys. Res.*, 93,1509-1527, 1988b.
- [14]. Cros, B.R., Delmas . D, Nganga, B., Clairac and J. Fontan (1988) Seasonal trends of ozone in Equatorial Africa: experimental evidence of photochemical formation, *J. Geophys. Res.*, 93, 8355-8366.
- [15]. Fontan et al 1992 J Fontan, A Minga, A Lopez Druilhet vertical ozone profile in a pine Forest Atmos
- [16]. Kundu, N. and M, Jain (1993): Total ozone trends over low latitude Indian stations, *Geophys. Res. Lett.*, 20,2881-2883.
- [17]. Randriambelo, T., Baray, J. L., Baldy, S., and Bremaud, P.: A Case study of extreme tropospheric ozone contamination in the tropics using in situ, satellite and meteorological data, *Geophys. Res. Lett.*, 26, 12 1287-1290, 1999.
- [18]. WOUDC (<http://www.lor.ec.gc.ca/woudc>)
- [19]. Website: [www.boh.ors](http://www.boh.ors), date accessed 4/2016.
- [20]. Kundu N., A.K. Saha (1987) Tropospheric ozone profiles near equator and meteorological parameters, *J Atmos- Chem*; 51243-253
- [21]. Kalicharran, S., R.D. Diab and F. Sokolic (1993): Trends in total ozone over South African stations between 1979-1991, *Geophys. Res. Lett* 20, 2877-2880.

- [22]. Masci, F. (1999); Algorithms for the inversion of lidar signals: Mie measurements in the stratosphere. *Ann Geophys.*, 42 (1), 71-83.
- [23]. Website: (GRADS) <http://www.grads.iges.org>, date accessed 4/2016.
- [24]. Website: NASA TOMS web page "<http://jwocky.gsfc.nasa.gov/>", date accessed 4/2016.
- [25]. Winterrath, T. et al., Enhanced O<sub>3</sub> and NO<sub>2</sub> in Thunderstorm Clouds: Convection or Production?, *Geophys. Res. Lett.*, 26, (9), 1291, 1294.
- [26]. Halocarbons and Other Gases. Emissions of Greenhouse Gases in the United States 1996. Energy Information Administration. 1997. Retrieved 2008-06-24.
- [27]. Stratospheric Ozone and Surface Ultraviolet Radiation. Scientific Assessment of Ozone Depletion: 2010 (PDF). WMO. 2011. Retrieved March 14, 2015.
- [28]. Farman, J.C; Gardiner, B.G.; Shanklin, J.D. (1985). Large losses of total ozone in Antarctica reveal seasonal ClO<sub>x</sub> / NO<sub>x</sub> interaction. *Nature* 335 (6016): 207-210. Biocode: 1985Natur.315.207 doi: 10.1038/315207a0.
- [29]. Grist and Nicholson, 2001 J.P. Grist, E. Nicholson A study of the dynamic factors influencing the rainfall variability in the West African Sahel *Journal of Climate*, 14 (2001), pp. 1337–1359.
- [30]. Grist et al., 2002 J.P. Grist, S.E. Nicholson, A.I. Barcion Easterly waves over Africa. Part II: observed and modeled contrasts between wet and dry years *Monthly Weather Review*, 130 (2002), pp. 212- 225. Full Text via Cross Ref/ View Record in Scopus/ citing articles (33).
- [31]. Omotosho, 2008 J.B. Omotosho Pre-rainy season moisture build-up and storm precipitation delivery in the West African Sahel *International Journal of Climatology*, 28 (2008), pp. 937-946. Full text via Cross Ref/View Record in Scopus/ Citing articles (7)
- [32]. Omotosho, 1988 J.B. Omotosho Spatial variation of rainfall in Nigeria during the 'little dry season' *Atmospheric Research*, 22 (1988), pp. 137–147 Article/PDF (447)/View Record in Scopus/ Citing articles (12).
- [33]. Lindzen and Holton, 1968 R.S. Lindzen, J.R. Holton A theory of the quasi-biennial oscillation *Journal of the Atmospheric Sciences*, 25 (1968), pp. 1095–1107 Full Text via Cross Ref/ View Record in Scopus/Citing articles (1).
- [34]. Dunkerton, 2001 T.J. Dunkerton Quasi-biennial and subbiennial variations of stratospheric trace constituents derived from HALOE observations *Journal of the Atmospheric Sciences*, 58 (2001), pp. 7-25 Full Text via Cross Ref/ View Record in Scopus/Citing articles (45).

- [35]. Dunkerton and Delisi, 1985 T.J. Dunkerton, D.P. Delisi Climatology of the equatorial lower stratosphere *Journal of the Atmospheric Sciences*, 42 (1985), pp. 376-396 Full Text via Cross Ref/View Record in Scopus/Citing articles (93).
- [36]. Baldwin et al., 2001 M.P. Baldwin, L.J. Gray, T.J. Dunkerton, K. Hamilton, P.H. Haynes, W.J. Randel, J.R. Holton, M.J. Alexander, I. Hirota, T. Horinouchi, D.B.A. Jones, J.S. Kinnerson, C. Marquardt, K. Sato, M. Takahashi The quasi-biennial oscillation *Reviews of Geophysics*, 39 (2001), pp. 179– 229. Full Text via Cross Ref/ View Record in Scopus/ Citing articles (743).
- [37]. Lindzen and Holton, 1968 R.S. Lindzen, J.R. Holton A theory of the quasi-biennial oscillation *Journal of the Atmospheric Sciences*, 25 (1968), pp. 1095–1107. Full Text via Cross Ref/ View Record in Scopus/Citing articles (1)
- [38]. Schönbein, 1840 C.F. Schönbein On the odour accompanying electricity and on the probability of its dependence on the presence of a new substance *Philos. Mag.*, 17 (1840), pp. 293-294. View Record in Scopus /Citing articles (10).
- [39]. Stratton, 1946 F.J.M. Stratton Prof. Charles Fabry, Foreign Member of the Royal Society *Nature*, 157 (1946), p. 362 doi: 10.1038/157362a0. Full Text via Cross Ref/ View Record in Scopus/Citing articles (1)
- [40]. De la Casinière, 2003 A. De la Casinière *Le rayonnement solaire dans l'environnement terrestre*, Éditions Publibook, Paris (2003), p. 264.
- [41]. Chapman, 1930 S. Chapman A theory of upper-atmosphere ozone *Mem. R. Meteorol. Soc*, 3 (1930), pp. 103- 125 View Record in Scopus/Citing articles (1).
- [42]. Molina and Rowland, 1974 M.J. Molina, F.S. Rowland Stratospheric sink for chlorofluoromethanes: chlorine atom-catalyzed destruction of ozone *Nature*, 249 (1974), pp. 810-812 doi: 10.1038/249810a0 Full Text via Cross Ref/ View Record in Scopus/ Citing articles (2146).
- [43]. Newchurch et al., 2003 J. Newchurch, E.-S. Yang, D.M. Cunnold, G.C. Reinsel, J.M. Zawodny, J.M. Russell III Evidence for slowdown in stratospheric ozone loss: First stage of ozone recovery *J. Geophys. Res.*, 108 (2003) doi: 10.1029/2003JD003471
- [44]. Yulaeva E, Holton JR, Wallace JM (1994). On the Cause of Annual Cycle in Tropical lower Stratospheric Temperature. *J. Atmos. Sci.* 51: 169-174.
- [45]. Rosenlof KH (1995). Seasonal Cycle of the Residual Mean Meridional Circulation in the Stratosphere, *J. Geophys. Res.* 100:5173-5191.
- [46]. Holton JR, Haynes PH, McInnes ME, Douglas AR, Rood RB, Pfister L (1995). Stratosphere

Troposphere Exchange. *Review of Geophy.* 33(4): 403-439.

- [47]. Zerefos CS, Tourpali K, Bojkov BR, Balis DS, Rognerund B, Isaksen ISA (1997). Solar activity total column ozone relationships: Observations and model studies with heterogeneous chemistry. *J. Geophys. Res.* 102: 1561-1569.